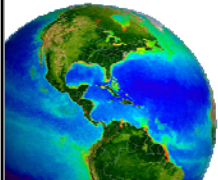


NASA vicarious calibration of on-orbit ocean color satellite instruments

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Darmstadt, Germany
7 May 2013



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presentation outline

- legacy of the NASA operational approach
- general method
- calibration of NIR (& SWIR) bands
- calibration of visible bands
- previous evaluations of alternative methods

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legacy of the NASA operational approach

H.R. Gordon, "In-orbit calibration strategy for ocean color sensors," Rem. Sens. Environ. 63, 265-278 (1998)

D.K. Clark et al., J. Geophys. Res. 102, 17209-17217 (1997)
D.K. Clark et al., NASA Tech Memo. 2004-211632 (2003)

R.E. Eplee Jr. et al., "Calibration of SeaWiFS. II. Vicarious techniques," Appl. Opt. 40, 6701-6718 (2001)

B.A. Franz, et al., "Sensor-independent approach to the vicarious calibration of satellite ocean color radiometry," Appl. Opt. 46, 5068-5082 (2007).

P.J. Werdell et al., Appl. Opt. 46, 5649-5665 (2007)
S.W. Bailey et al., Appl. Opt. 47, 2035-2045 (2008)

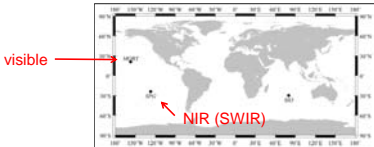
In place since SeaWiFS Reprocessing 4 (R2002)

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general method

the vicarious calibration:

- minimizes the difference between satellite & ground truth $L_w(\lambda)$
- modifies the integrated instrument + atmospheric correction system
- uses a single set of fractional gains (unity = no correction)
- assumes temporal trends are independently removed
- currently uses the full spectral response of each sensor band



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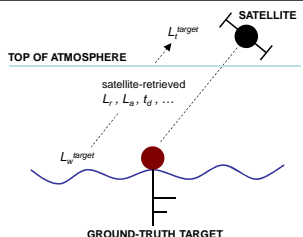
general method

the water-leaving signal from target is:

- adjusted to solar/sensor geometries of satellite overpass (using BRDF correction of Morel et al. 2002)
- propagated to top-of-atmosphere using satellite-retrieved aerosol properties

vicarious gains are calculated as:

- $g(\lambda) = L_{t, target}(\lambda) / L_{t, satellite}(\lambda)$
- $g(\lambda) = \sum g_i(\lambda) / N$ (specifically, mean of SIQR)



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vicarious calibration of NIR bands

$$L_t(NIR) = L_{r, g/w/c} \text{ (calculated) } (NIR) + L_w(NIR) + L_w(NIR) \text{ (0)}$$

requires two assumptions:

- L_w in two NIR bands negligible (= 0)
- calibration of longest NIR band is perfect (e.g., $g(865) = 1$ for SeaWiFS)

calibration of remaining NIR band (e.g., 765 for SeaWiFS):

- possible when aerosol type is known; the associated model can be used in combination with $L_w(865)$ to predict $L_w(765)$
- operationally executed using a 15x15 pixel target in the South Pacific Gyre (aerosol model r70f10v01; $\alpha = 0.685$; assigned using Tahiti AERONET site)
- remains spatially/temporally independent of visible band calibration

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vicarious calibration of visible bands

$$L_v(\text{VIS}) = L_{\text{gr}}(\text{VIS}) + L_a(\text{VIS}) + L_w(\text{VIS})$$

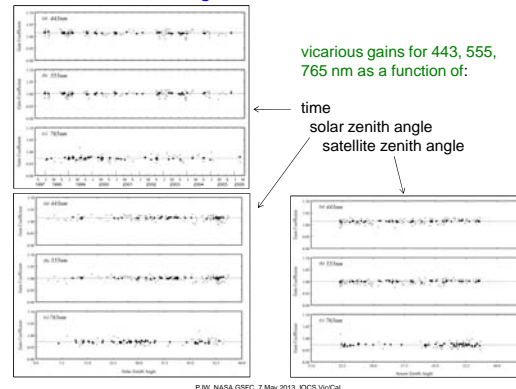
← calculated
← from ground-truth target

calibration of visible bands:

- uses calibrated NIR bands to dynamically determine local aerosol type & concentration at ground-truth target site, which provides $L_a(\text{VIS})$
- operationally executed using $L_w(\text{VIS})$ from the Marine Optical Buoy (MOBY) & a 5x5 pixel target centered on the buoy location
- exclusion criteria follows S.W. Bailey and P.J. Werdell, "A multi-sensor approach for the on-orbit validation of ocean color satellite data products," Rem. Sens. Environ. 102, 12-23 (2006)

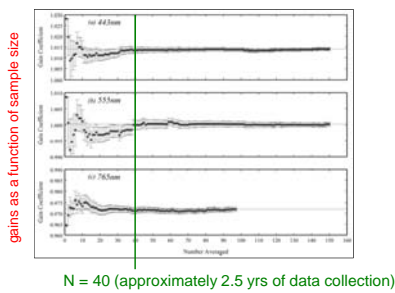
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vicarious gain series for SeaWiFS



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vicarious gain series for SeaWiFS



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previously tested alternative approaches

methods for gain estimation:

- slope & offset in lieu of ratio (~ slope with offset = 0)
- gains calculated at L_w in lieu of L_a
- simultaneous use of *in situ* aerosol optical thickness
- zero Rayleigh (courtesy Zia Ahmad)
- zero aerosol (no NIR contributions by aerosols or water)
- NIR -> SWIR; SWIR -> NIR

data sources:

- modeled $L_w(\lambda)$ (Werdell et al. 2007)
- alternative *in situ* data (NOMAD, BOUSSOLE; Bailey et al. 2008)

sensitivity analyses (Franz et al. 2007):

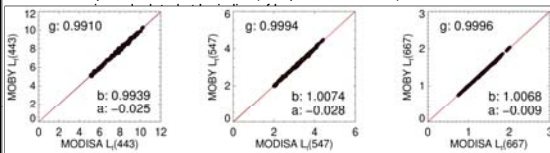
- NIR-only calibration (no visible vicarious calibration)
- $g(865) = 1.04$ or 0.96 (+/- 4% from unity)
- alternative aerosol models

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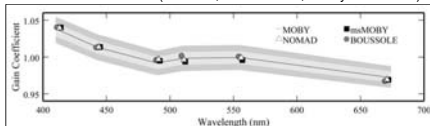
previously tested alternative approaches

methods for gain estimation:

- slope & offset in lieu of ratio (~ slope with offset = 0)



- alternative *in situ* data (NOMAD, BOUSSOLE; Bailey et al. 2008)



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acknowledgements & contact info

additional vicarious calibration info available from:

- Sean Bailey
- Bryan Franz
- Zia Ahmad
- Chris Proctor
- Gene Eplee
- Bob Barnes

<http://oceancolor.gsfc.nasa.gov>

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alternative *in situ* data (Bailey et al. 2008)

Table 2. Vicarious Gain Coefficients for Standard Method*

Source	N	412	443	490	510	555	670
MOBY	166	1.0368	1.0132	0.9918	0.9982	0.9993	0.9729
(σ)		(0.0066)	(0.0076)	(0.0080)	(0.0080)	(0.0089)	(0.0077)
CV	1796	1.777	1.413	1.403	1.403	1.401	1.439
NOMAD	64	1.0365	1.0135	0.9967	0.9962	0.9969	0.9683
(σ)		(0.0113)	(0.0113)	(0.014)	(0.017)	(0.0113)	(0.009)
LTPD	5	0.5207	0.01490	0.2484	-0.1003	-0.0200	-0.1854
BOUSSOLE*	46	1.6827	1.0129	0.9961	1.0013	1.0007	0.9872
(σ)		(0.065)	(0.027)	(0.032)	(0.031)	(0.021)	(0.006)
LTPD		0.4827	-0.0144	0.1143	0.1653	0.0790	-0.3908

*Gain coefficients using the threshold criteria defined in 10. The standard deviations are shown in parentheses. CV, threshold increased to 0.25 mg m⁻³ for the BOUSSOLE data set to bring the N to a minimum of 40. The 412 nm data for BOUSSOLE used only 9 points.

Table 3. Vicarious Gain Coefficients Using moMOBY Data*

	412	443	490	510	555	670
MOBY	1.0401	1.0136	0.9949	0.9927	0.9954	0.9691
RPD	0.3149	0.0595	0.3126	-0.4508	-0.3502	-0.3908

*Gain coefficients for MOBY data processed to mimic a multi-spectral, COTS instrument.

Table 4. Validation Results*

Band	N	MOBY		moMOBY		NOMAD/BOUSSOLE	
		Ratio	% DME	Ratio	% DME	Ratio	% DME
Z ₄₉₀ (412)	154	1.005	11.762	1.005	11.83	0.814	0.997
Z ₄₉₀ (443)	296	0.938	15.96	0.938	16.10	0.824	0.822
Z ₄₉₀ (490)	296	0.918	13.62	0.909	12.77	0.706	0.803
Z ₄₉₀ (510)	127	0.953	11.97	0.948	12.01	0.815	0.985
Z ₄₉₀ (555)	296	0.961	15.95	0.950	17.45	1.223	0.868
Z ₆₇₀ (570)	253	1.719	87.21	1.218	81.18	12.09	1.091
C _{oc2}	262	1.051	27.40	0.998	27.80	1.269	1.050

*Deep-water validation results for satellite data processed with the standard MOBY-derived g_r , the moMOBY-derived g_r , and the weighted average of the NOMAD- and BOUSSOLE-derived g_r . The ratio column shows the median ratio of the satellite to *in situ* measured values, the %DME column shows the median percent difference for the same, and the Abs. LTPD column shows the absolute estimated percent difference between the satellite values estimated using the MOBY-derived g_r and the satellite values estimated from the alternative sea truth data derived g_r .

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alternative *in situ* data (Bailey et al. 2008)

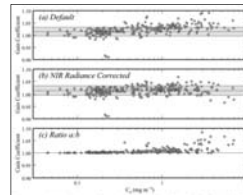


Fig. 4. $g_r(443 \text{ nm})$ as a function of satellite-derived C_c for (a) NIR-uncorrected data set, (b) the NIR-corrected data set and the ratio of the two, and (c) ratio of (a) to (b).

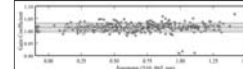


Fig. 6. $g_r(443 \text{ nm})$ derived from MOBY (open circles), as well as the NIR corrected $g_r(443 \text{ nm})$ derived from NOMAD and BOUSSOLE data sets (filled circles) versus the satellite-estimated Algorithm (510 - 865 nm) exponent.

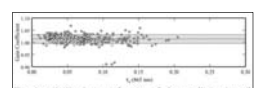


Fig. 5. $g_r(443 \text{ nm})$ as a function of the satellite-estimated $C_c(665 \text{ nm})$. Data from MOBY (open circles), as well as the NIR-corrected NOMAD and BOUSSOLE data sets (filled circles) are shown.

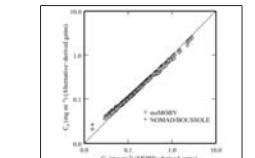


Fig. 7. Satellite-derived chlorophyll estimated from the two alternative g_r gain sets (moMOBY and NOMAD/BOUSSOLE) plotted versus the corresponding chlorophyll estimated from the standard MOBY g_r .

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