Atmospheric trace gas (NO₂ and ozone) dynamics and impacts on Ocean Color retrievals

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IOCS 2019 – Atmospheric correction under complex or extreme conditions/environments

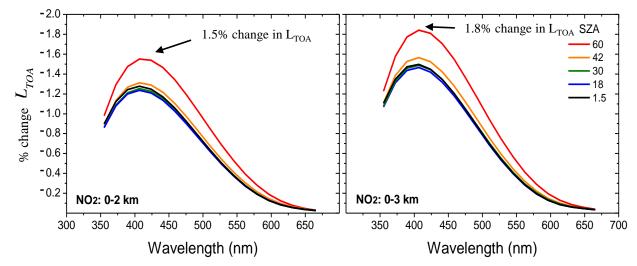
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VC NO₂ [molec cm⁻²] 4.0 10¹⁶ 2.0 10¹⁶ 1.0 10¹⁶ 5.0 10¹⁵ 2.0 10¹⁵ 1.0 10¹⁵ 5.0 10¹⁴

NASA

55

Percent change in TOA signal, per 1 DU change in atmospheric NO_2

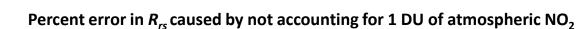


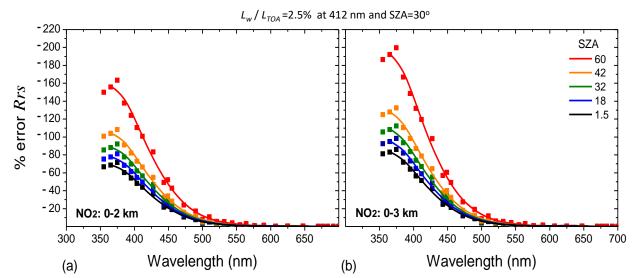
The impact on L_{TOA}

- has a strong spectral dependence: max in 400-420 nm, due to spectral shape in NO₂ abs. cross sections
- has a SZA dependence: because of the larger slant path with increasing SZA
- depends on NO₂ vertical distribution, and becomes larger as the NO₂ is distributed at higher altitudes

Tzortziou et al. 2014, Geophys. Res. Oceans, 119, 3834–3854, doi:10.1002/2014JC009803







The impact on R_{rs}

- has a strong spectral dependence: max in 350-400 nm, due to spectral dependence of Lw/L_{TOA}
- has a SZA dependence: because of the larger slant path with increasing SZA
- depends on NO₂ vertical distribution, and becomes larger as the NO₂ is distributed at higher altitudes

Tzortziou et al. 2014, Geophys. Res. Oceans, 119, 3834–3854, doi:10.1002/2014JC009803





May-June 2016

3 aircraft, 2 research vessels > 350 scientists and other personnel in the field



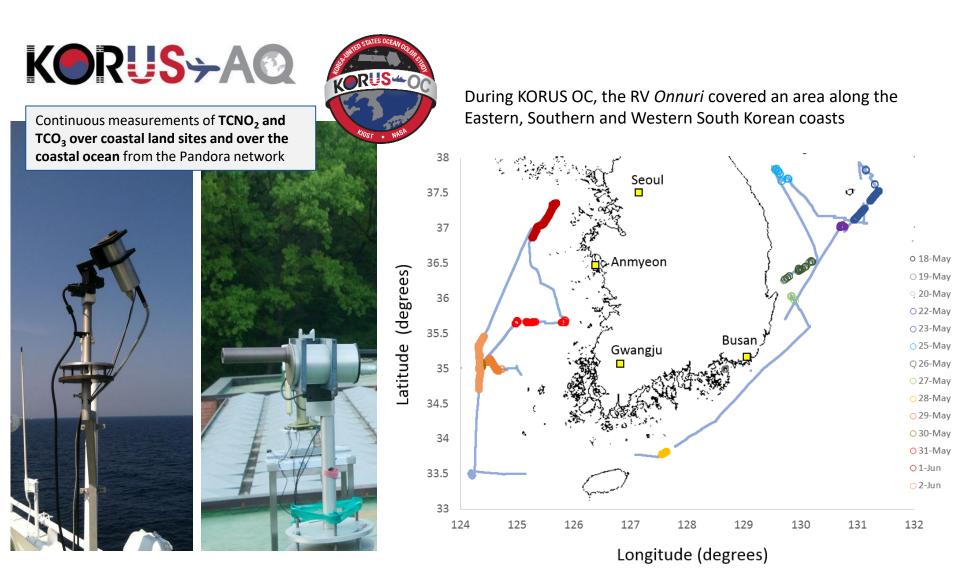
Objectives:

- Capture spatial dynamics and diurnal variability in NO₂ and O₃ using shipboard and satellite observations
- Examine differences in air quality across the land-ocean interface
- Assess source contributions to atmospheric pollution over these coastal waters



3 airplanes with 37 different instruments onboard (MOS+P, GeoTASO)





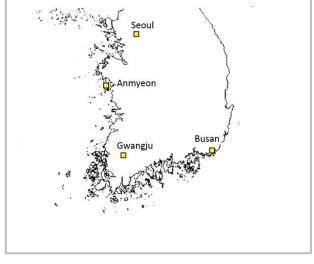


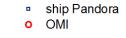


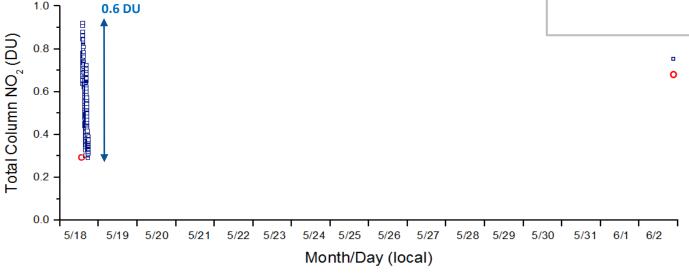
TC NO₂ varied from 0.07 DU to 0.92 DU during KORUS OC

> 35 km distance from Busan

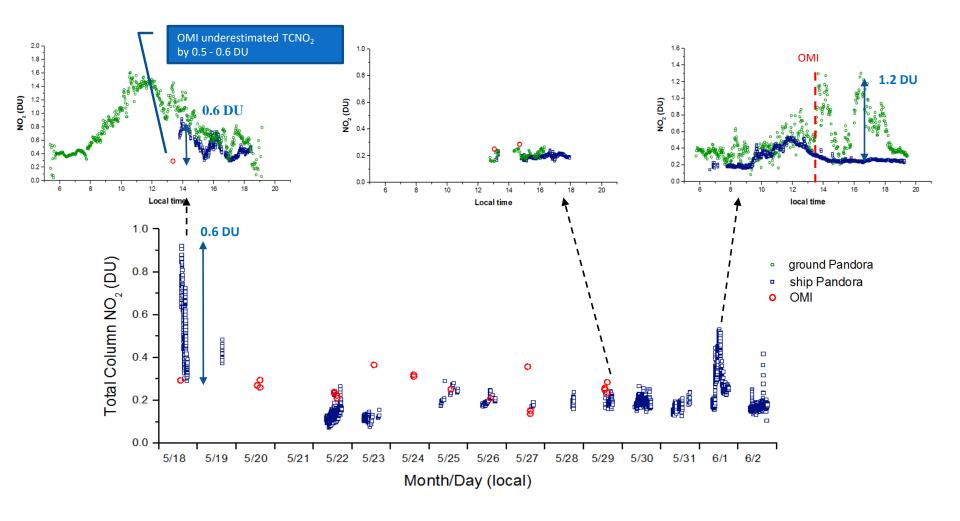
- OMI underestimated TCNO₂ in areas of high NO₂ pollution, and overestimated NO₂ over the ocean under relatively clean-air conditions
- Change as large as 0.6 DU within 3 hours







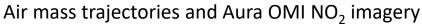


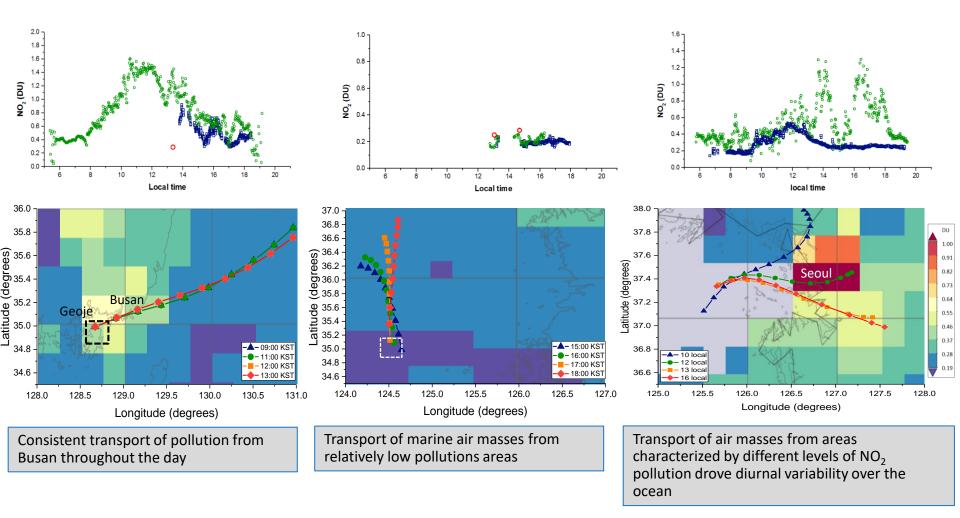






Backward air parcel trajectories simulated using HYSPLIT4 to determine the origin of air masses over *RV Onnuri*

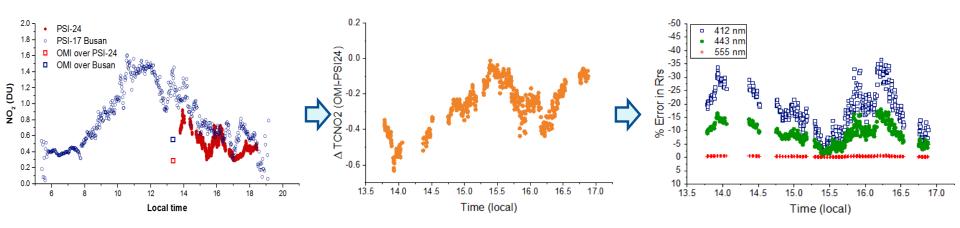








What would be the impact of using daily **TCNO**₂ values from OMI (GOME and SCIAMACHY are much coarser resolution) in atmospheric correction algorithms for ocean color products in these coastal waters ?



- TCNO₂ uncertainty of 0.6 DU during the day
- Some of the largest differences occurring at 14:00 and 16:00 local time corresponding to 25° and 50° SZA

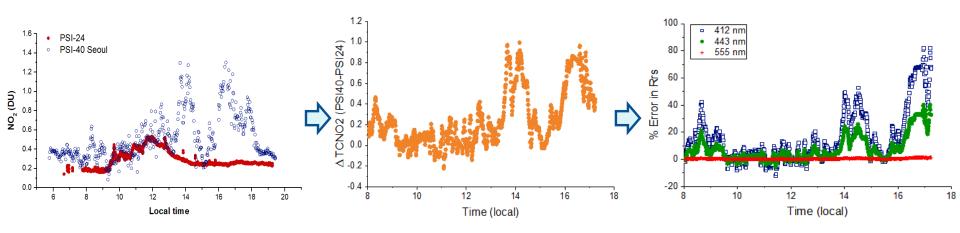
Impact:

- a false change in R_{rs}(412) by 37% during the day
- Smaller error for 443 nm, and almost negligible error at 555 nm, introducing an uncertainty not only in the absolute amount of R_{rs}, but also in its spectral shape





In the absence of information on the diurnal variability of TCNO₂ over the ocean (from shipboard or satellite platforms), the diurnal variability in TCNO₂ measured by **ground-based instruments at nearby coastal land sites** could be used as a reasonable alternative.



- TCNO₂ uncertainty as large as 1 DU during the day
- Some of the largest differences occurring in the afternoon corresponding to large SZAs

Impact:

- a false change in R_{rs}(412) by 43% in the morning and as large as 82% in the afternoon
- Spectral dependence on error in R_{rs} introduced an uncertainty not only in the absolute amount of R_{rs} but also in its spectral shape





- Small scale variability in TCNO₂, needs to be taken into consideration when attempting to retrieve short-term coastal ocean processes from satellite ocean color imagery.
- Need for improvements of satellite retrievals of atmospheric NO₂ over the ocean. Uncertainties in the OMI retrievals (e.g., NO₂ profile shape and surface reflectivity) contributed to the discrepancies between OMI and Pandora. Current research is focusing on reducing these uncertainties over land and over the ocean.
- Need for high-spatial resolution, high-temporal resolution observations of atmospheric NO₂ from shipboard platforms and integration of these measurements to atmospheric correction approaches
- Need for high-spatial and high-temporal resolution observations of atmospheric NO₂ from space (e.g., TEMPO, TROPOMI, GEMS, Sentinel-4, Sentinel-5)



Discussion: recommendations for AC over complex atmosphere



1. Advancement in algorithms with current sensors:

Is rigorous modeling of the aerosol required (IOP, PF, ALH....)?



How global assimilated aerosol transport models improves algorithms? Is there a need for high resolution models (~ few kilometers)?



2. Advancement in observation platforms:

How can UV improve the aerosol correction? What are the technological limitations in designing a high radiometric quality UV spectrometer with high spatial resolution?



- Can hyperspectral observations:
- of the O2 bands improve layer height detection (limitations to high optical depth)?
- of the blue bands (400-450 nm) be used to correct for NO2?



Is there a necessity for a LIDAR/polarimeter/multiangular instrument combined with an ocean color sensor for aerosol vertical column profile detection?

